### EXTRATERRESTRIAL ATMOSPHERIC PHOTOCHEMISTRY

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# **1. Introduction**

Our knowledge of the atmospheres of the planets and satellites in the solar system has advanced enormously [1 - 3] in the years since the first issue of the *Journal of Photochemistry* appeared. Orbiters, "fly-bys" and landers have reached the atmospheres of distant bodies to provide the first *in situ* measurements, and these data have been supplemented by the results from increasingly sophisticated Earth-based investigations. Table 1 summarizes some of the most important recent space missions that have included atmospheric observations as an objective [4 - 19], and the table also shows what is planned for the immediate future.

In the paragraphs that follow I have presented a very personal selection of topics that are specifically photochemical in nature.

#### **2.** Inner planets

One of the fascinating aspects of the photochemistry of the inner planets has been the identification of the part that catalytic cycles play in determining composition [20 - 23]. For Mars and Venus, such cycles are vital, because photolysis of the bulk atmospheric constituent,  $CO_2$ ,

$$CO_2 + h\nu(\lambda \le 204 \text{ nm}) \longrightarrow CO + O$$
 (1)

produces CO (and O) which could not otherwise be oxidized back to  $CO_2$ . Present day CO concentrations could be produced in a few years on Mars and a few hundred years on Venus, and the entire atmospheres would be modified in a few thousand years. On Mars [20, 21], catalytic cycles involving HO<sub>x</sub> species can be readily identified, with the reaction

$$CO + OH \longrightarrow CO_2 + H$$
 (2)

playing a dominant role in the oxidation of CO. Odd hydrogen compounds are supplied by photochemical decomposition of  $H_2O$ , either directly by photolysis or in attack by  $O(^1D)$  derived from  $CO_2$ ,  $O_2$  and  $O_3$  photolysis. A small branch of the  $HO_x$  reactions leads to the formation of  $H_2$  that ultimately escapes as H from the Martian exosphere. Since the reaction is fed

Date	Name	Objective	References
July - September 1976	Viking 1, 2	Mars Entry Science and Lander	[4, 5]
December 1978	Pioneer Venus	Orbiter and "Bus" carrying entry probes to Venus	[6 - <b>9]</b>
December 1978	Venera 11, 12	Soviet probes to Venus	[9-11]
March 1979	Vovager 1	Closest approach to Jupiter	[12]
July 1979	Vovager 2	Closest approach to Jupiter	[13]
September 1979	Pioneer 11	First ever fly-by of Saturn	[14, 15]
November 1980	Vovager 1	Closest approach to Saturn	[16 - 18]
August 1981	Vovager 2	Closest approach to Saturn	[17 19]
March 1982	Venera 13, 14	Soviet sounders to Venus	[11, 10]
June 1984	"VEGA" (Venera 15, 16)	Launch December 1984 to visit Venus (June 1984) and on to comet Halley (April 1986) (U.S.S.BFrance)	
April 1986	Giotto	Comet Halley mission (ESA; launch, summer 1985)	
March - April 1986	Planet C	Comet Halley mission (Japan)	
1986	Voyager 2	First ever fly-by of Uranus	
1989	Galileo	Jupiter orbiter (NASA: launch, 1986)	
1989	Voyager 2	First ever fly-by of Neptune	

TABLE 1

Some recent and proposed missions

ESA, European Space Agency; NASA, National Aeronautics and Space Administration.

by the leak of  $H_2O$  that is permanently destroyed, escape of two hydrogen atoms is necessarily accompanied by the escape of one oxygen atom [21]. Ionospheric processes [24] are responsible for this O loss, as well as for the loss of N [21]. The  $H_2O$  lost over the age of the solar system would correspond to an ice layer of 2.5 m thickness. Incidentally, since there is only an *atmospheric* reservoir of nitrogen, the slightly higher loss rate of <sup>14</sup>N compared with <sup>15</sup>N has led to an isotopic enrichment of the heavier isotope in the atmosphere.

The catalytic cycles which effect the union of CO and O on Venus are rather more complex and subtle than those operating on Mars, because there is insufficient  $H_2O$  to yield the necessary concentration of  $HO_x$  radicals [23]. Considerable interest attaches to the possible roles [22, 23] of  $ClO_x$  and  $SO_x$ species in catalytic cycles because of the presence of HCl and  $SO_2$  in the upper atmosphere. In the Earth's atmosphere,  $ClO_x$  only catalyses the conversion of odd oxygen to  $O_2$ , but on Venus a dual role is possible, with two sequences leading to recombination of O with CO. Three-body formation of chloroformyl radicals (ClCO) and then of peroxychloroformyl radicals (ClCO<sub>3</sub>) is involved in both cases:  $Cl + CO + M \longrightarrow ClCO + M$  (3)

$$ClCO + O_2 + M \longrightarrow ClCO_3 + M$$
 (4)

followed by attack of either Cl or O on  $ClCO_3$  to yield  $CO_2$ . Sulphurcontaining species are also implicated in the (photo)dissociation of  $O_2$  [23], with sulphur atoms, ultimately derived photochemically from  $SO_2$ , attacking the molecular oxygen. More interestingly, the SO product of  $SO_2$  photolysis can react with ClO:

$$SO + ClO \longrightarrow Cl + SO_2$$
 (5)

to synergize the  $ClO_x$  catalytic chain. The mechanism bears an interesting analogy with the tropospheric oxidation cycle on Earth, of which a part may be represented

$$HO_2 + NO \longrightarrow NO_2 + OH$$
 (6)

$$NO_2 + h\nu \longrightarrow NO + O$$
 (7)

On Venus, SO<sub>2</sub> plays the role of NO<sub>2</sub>, and  $ClO_x$  the role of HO<sub>x</sub>. Mesospheric hazes of ice, dust and acid aerosols [25] bear a further analogy with terrestrial phenomena!

Electronically excited species do not appear to play any major direct part in the bulk chemistry of the neutral atmospheric components, although they may be involved in ionospheric processes [24, 26] that ultimately affect the neutral chemistry. They are also, of course, intimately connected with airglow emissions. UV airglow studies [27 - 30] have afforded considerable insight into the distribution of atomic and other minor constituents of the atmospheres of Mars and Venus, but we concentrate here on emissions from the metastable singlet states of  $O_2$  [31]. IR Atmospheric Band emission, due to the  $O_2({}^1\Delta_g \rightarrow {}^3\Sigma_g^-)$  transition, is observed in the airglow of both Mars [32, 33] and Venus [34]. In the Martian atmosphere,  $O_2({}^1\Delta_{\sigma})$  is almost certainly excited, as it is in the terrestrial dayglow, by the UV photolysis of  $O_3$  [31]; seasonal and latitudinal intensity variations [33] are entirely consistent with the variability of  $[O_3]$ .  $O_3$  photolysis can account [35] for only about 10% of the observed intensity on Venus because of the low oxygen column densities [36]. Alternative excitation mechanisms that have been proposed include [31] recombination of atomic oxygen and exothermic reactions of  $ClO_x$  species, but the proposals await experimental confirmation. A big impetus to the study of the higher lying (about 4.05 eV)  $c^{1}\Sigma_{u}^{-}$  state of O<sub>2</sub> was provided by the identification [37] of a Herzberg II ( $c \rightarrow X$ ) progression in a nightglow spectrum obtained [38] by the Soviet Venera 9, 10 mission. Subsequently, it was shown [39] that the  $c \rightarrow X$  emission was also present in the Earth's nightglow, although much weaker [39, 40] than in the Venusian nightglow, and with an entirely different vibrational distribution. In the Martian airglow, the Herzberg II system is weaker [41] even than it is in the Earth's. Laboratory studies [37, 42 - 45] have shown that recombination of  $O(^{3}P)$  atoms populates  $O_2(c^{1}\Sigma_{\mu})$ , especially when  $CO_2$  is the third body.

Atomic oxygen is a particularly efficient [44, 46] quencher of  $O_2(c^1\Sigma_u^-)$ , and the relative intensities of the Herzberg II bands on Venus, Earth and Mars may reflect the abundance of O as well as of  $O_2$  and the nature of the third body [44, 47]. Great importance attaches to the question of whether  $O_2(c^1\Sigma_u^-)$  can excite  $O({}^1S)$ ,  $O_2(a^1\Delta_g)$  or  $O_2(b^1\Sigma_g^+)$  in collisions with O or  $O_2$  [31, 44 - 49], and it seems that these studies will continue to excite much interest.

We can hardly leave the topic of non-equilibrium emission processes without some reference to "natural lasers" [50] operating in the CO<sub>2</sub> atmospheres of Venus and Mars. Absorption of energetic solar radiation by CO<sub>2</sub>, followed by *partial* collisional deactivation, may populate vibrationally excited CO<sub>2</sub> in the upper atmospheres [51]. Thermally emitted photons from the lower atmosphere can then stimulate emission from these molecules. Stimulated emission of this kind may account for up to 7% of the total radiation at  $\lambda \approx 10 \,\mu\text{m}$  [51], but speculations are now appearing [50, 51] that this weak natural laser could be turned into a huge man-made one by placing large mirrors in orbit parallel to one another!

## 3. Outer planets and Titan

Jupiter and Saturn are made up *almost* entirely of hydrogen and helium, although  $CH_4$  and  $NH_3$  are present in abundances rather greater than the solar ones; Uranus and Neptune probably have a rocky core, and a correspondingly greater departure from solar composition [52, 53]. The temperature structures in the atmospheres of the four planets, and of Saturn's moon Titan, all show minima, with a tropopause separating regions analogous to Earth's troposphere and stratosphere. The temperature minima act as "cold-traps" which limit the mixing ratios of condensable gases in the upper atmosphere, with condensates remaining as aerosols or clouds at the appropriate levels.

Above the clouds on the planets, photochemical transformations can take place [54]. Molecular hydrogen is converted, directly or via an ionic mechanism, to hydrogen atoms at  $\lambda < 100$  nm. CH<sub>4</sub> is photolysed at longer wavelengths to yield mainly singlet and triplet  $CH_2$  radicals [54]. Hydrogen abstraction by  $CH_2$  leads to  $CH_3$  and ultimately  $C_2H_6$ .  $C_2H_4$  is also probably formed, but is photolysed itself to yield  $C_2H_2$  which is photochemically rather stable (because the dissociation products  $C_2H$  and  $C_2$  react with  $H_2$  to regenerate  $C_2H_2$ ). Higher hydrocarbons, and possibly even polymeric materials, can be formed by reactions of  $C_2H_2$  (as, for example, with <sup>3</sup>CH<sub>2</sub> to yield methyl acetylene,  $CH_3CCH$ , an observed species).  $NH_3$  is photolysed to  $NH_2$ , some of which combines to yield hydrazine,  $N_2H_4$ , as a condensed haze. A minor photochemical pathway leads, via  $N_2H_4$ , to  $N_2$ , and, were it not for the condensation of  $N_2H_4$ , molecular nitrogen would be the major nitrogen-containing species [54]. Analogous reactions lead from PH<sub>3</sub> to  $P_2H_4$ , with some elemental phosphorus being formed [54, 55]. There is naturally some speculation [55] about whether phosphorus contributes to

the wide variety of colours found on the bodies of the outer solar system. Colouration by complex organic molecules is another possibility, and in this context the presence of HCN, at least on Jupiter, may be significant. It is, in fact, rather difficult to account for the formation of HCN in these atmospheres [56]; one promising pathway is the photolysis of aziridine (ethyleneimine), itself formed directly from NH<sub>2</sub> radicals and  $C_2H_2$  [56].

Although the atmospheres of Jupiter and Saturn are usually thought of as totally reducing, the discovery of CO on Jupiter proves that photochemical processes involving oxygen species do occur. An extraplanetary source of oxygen seems implicated [57],  $H_2O$  being supplied by meteors, material from the satellites and even from the rings for Saturn. Oxygen in any form is virtually bound to be converted to CO, which is then chemically and photochemically stable.

Saturn's largest satellite, Titan, is the only satellite in the solar system to possess a massive atmosphere [58 - 60]. Emission lines from molecular and atomic nitrogen, characteristic of electron-excited  $N_2$ , are emitted from the upper atmosphere. The atmosphere is predominantly  $N_2$  (with a few per cent of CH<sub>4</sub>, and some other minor compounds) exerting a surface pressure of 1.6 (Earth) atm. Titan is the only body in the solar system besides Earth that has an atmosphere composed largely of nitrogen. The satellite is covered by coloured clouds that must be aerosols derived from gaseous organic compounds, and are thus Titan's equivalent of photochemical smog; above the cloud layer lies a thinner haze of aerosol particles [61].

Escape of H and  $H_2$  from Titan's atmosphere is ensured by the combination of very low escape velocity  $(2.1 \text{ km s}^{-1})$  and the thermal structure of the atmosphere. Build-up of heavier hydrocarbons, such as  $C_2H_6$  and  $C_3H_8$ , occurs at the expense of  $CH_4$ , and  $H_2$  is a minor constituent. Absence of  $H_2$ , and the presence of N<sub>2</sub>, modify considerably Titan's atmospheric chemistry from that of Jupiter or Saturn [58, 59]. Radicals such as  $CH_3$  or  $NH_2$  no longer return most frequently to  $CH_4$  or  $NH_3$ . Hydrogen-deficient hydrocarbons are favoured, with mixing ratios of  $C_2H_6$  and  $C_2H_2$  far larger than on Saturn;  $C_2H_4$  is detected on the satellite but not on the planet. Nitrogen quenches  ${}^{1}CH_{2}$  to  ${}^{3}CH_{2}$ , and  $C_{2}H_{2}$  formation is particularly important. The radicals  $CH_3$ ,  ${}^{3}CH_2$  and  $C_2H$  can yield [59] many of the compounds, such as CHCCCH (diacetylene) and CH<sub>3</sub>CCH, observed [58] in the atmosphere. HCN formation is easier to explain on the satellite than on Saturn because the nitrogen source is  $N_2$  itself, with known routes to HCN existing from photochemically derived N and N<sup>+</sup>. Photolysis of HCN generates CN that can react with known atmospheric constituents to yield two other species detected in the atmosphere (HCCCN and  $C_2N_2$ ).

Recent discoveries of CO [62] and CO<sub>2</sub> [63] on Titan have opened up new vistas in the chemistry of reducing atmospheres. Continued input of  $H_2O$ to the satellite is demanded in order to maintain any CO<sub>2</sub> at all; without  $H_2O$ , photolysis would convert CO<sub>2</sub> to CO, and there would be no route for reversing the process in reaction (2), as we discussed in connection with Mars. If the large postulated influx of  $H_2O$  is borne out by further observation,  $H_2O$  could be an important species on Saturn as well. CO may therefore be present in Saturn's upper atmosphere as it is in Jupiter's. The chemistry of  $H_2O$ , CO and CO<sub>2</sub> may ultimately prove to be as significant to reducing atmospheres as that of the "trace constituent" reduced compounds (e.g.  $CH_4$ ,  $H_2$ ) is for oxidizing ones [63].

For the immediate future, refinement of the data from the Voyager encounters with Jupiter, Saturn and their satellites continues to yield new information which challenges the interpretative skills of photochemists. If all goes well, Voyager 2 will visit Uranus and Neptune in 1986 and 1989 and should provide a wealth of observational data. In 1986 as well, there is a chance of an encounter with Halley's comet to investigate the tail of gas that bears some similarity to an atmosphere. Knowledge of chemistry at the limits of the solar system should thus be expanding enormously when the *Journal of Photochemistry* reaches its fiftieth volume.

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